

# The Response of Atlantic Tropical Cyclones to Suppression of African Easterly Waves

Christina M. Patricola, Lawrence Berkeley National Laboratory R. Saravanan and Ping Chang, Texas A&M University

Geophysical Research Letters (2018)



PIRATA-23 – Marseille, France – October 23, 2018



- 1) warm upper-ocean temperature
- 2) unstable atmosphere
- 3) moist mid-troposphere
- 4) weak vertical wind shear
- 5) "seedling" low-pressure disturbance in Atlantic, usually AEWs

## African Easterly Waves (AEWs)

- trough of low pressure, 850 600 hPa
- generate over the Sahel with 2-10 day periodicity
- two preferred tracks: north and south of AEJ
- peak with the Atlantic hurricane season



(Gray 1968; Carlson 1969; Burpee 1974; Reed et al. 1977; Diedhiou et al. 1998; Chen et al. 2008)

## Synoptic: strongly linked

• 85% of major Atlantic hurricanes originate from AEWs

## Interannual: unknown

- Depends on reanalysis, time period, vertical level
- Covariability in TCs, AEWs, Atlantic SST, and West African monsoon
  - $\rightarrow$  causality is difficult to unravel

(Dunn 1940; Frank 1970; Landsea 1993; Russell et al. 2017; Thorncroft and Hodges 2001; Hopsch et al. 2007; Belanger et al. 2012; Landsea and Gray 1992; Grist 2002; Martin and Thorncroft 2014; Frank 1970; Dunkerton et al. 2009; Agudelo et al. 2011; Satoh et al. 2013)





## **ITCZ** wave instability

observed in Atlantic and Pacific

## **Disturbances from monsoon trough**

 observed in NW Pacific, where 10-25% of typhoons form from tropical waves

## Self-aggregation of convection

• in rotating radiative-convective equilibrium simulations

(Agee 1972; Thompson and Miller 1976; Kieu and Zhang 2008; Cao et al. 2013; Yokota et al. 2012; Bretherton et al. 2005; Nolan et al. 2007; Held and Zhao 2008; Khairoutdino and Emanuel 2013; Zhou et al. 2014; Reed and Chavas 2015; Wing et al. 2016; Frank 1988; Ritchie 1995; Lander 1994; Chen et al. 2004; Chen et al. 2008; Yoshida and Ishikawa 2013)





- Lack of clear seasonal connection between AEWs and Atlantic TCs
- Several alternative TC genesis mechanisms



Unclear whether changes in AEWs are a source of uncertainty in seasonal predictions and future projections of TC activity.



# **Regional climate model experiments**



#### Weather Research and Forecasting Model (WRF)

- Eastern lateral edge location: include TC development, exclude AEW genesis over Sahel
- LBCs: 6-hourly NCEP CFSR; SST: daily NOAA-OI V2
- 27 km resolution
- 10 ensemble members (1 Jun -1 Dec 2005)
- **Control: 2005** AEWs contributed to 75% of Atlantic TC genesis (Beven et al. 2008)
- AEW\_suppressed: 2-10 day variability removed from eastern LBC (5°S-30°N) w/ Lanczos filter



# **Downstream impact of AEW filtering**





- Filtering is largely effective at diminishing AEWs
- Model generates its own synoptic variability ~20° of longitude downstream from the lateral edge (no nudging applied within model domain)

Standard deviation of the 10-day high-pass filtered meridional wind  $(m s^{-1})$  at 15°N and 700 hPa. Wind data cover 1 Jul – 1 Oct, 2005.



A vorticity maximum enters via the eastern lateral edge and develops into a TC in temporally and spatially coherent manner in all control ensemble members.



control (8 ensemble members)

The sum of 6-hourly positive relative vorticity at 850 hPa ( $10^3$  s<sup>-1</sup>; shaded green) and daily snapshots of 10-m wind speed at 00z (m s<sup>-1</sup>; contour; color denotes day of Aug) over Aug 20–27, 2005.





Filtering removes the vorticity maximum present in the control  $\rightarrow$  TCs develop, but not in a temporally and spatially coherent manner.



The sum of 6-hourly positive relative vorticity at 850 hPa ( $10^3$  s<sup>-1</sup>; shaded green) and daily snapshots of 10-m wind speed at 00z (m s<sup>-1</sup>; contour; color denotes day of Aug) over Aug 20–27, 2005.



# Seasonal impact of AEW filtering



- Prescribing AEWs: spatial and temporal patterns in ensemble members resemble ensemble mean → AEWs initiated TC genesis in a coherent way throughout the season
- Filtering AEWs: little spatial and temporal coherence among ensemble members



Hovmöller diagrams of meridional wind (m s<sup>-1</sup>) at 700 hPa and averaged 14°N–16°N.

# Suppression of AEWs produced no significant change in seasonal Atlantic TC number

Supported by results from synthetic track models and statistical models

- Input: atmospheric thermodynamic state, vertical wind shear, SST
- No information about AEW variability is given to the model.
- Largely reproduce interannual variability of Atlantic TC activity

(e.g., Emanuel et al. 2008; Emanuel 2010; Saunders et al. 2017)



|  | control | AEW<br>suppressed | % change | p-value |
|--|---------|-------------------|----------|---------|
| Number of TCs/season                   | 19.5    | 20.2              | +4%      | 0.64    |
| Number TC days/season                  | 105     | 117               | +11%     | 0.17    |
| ACE (10 <sup>4</sup> kt <sup>2</sup> ) | 168     | 192               | +15%     | 0.07    |







### Regional model experiments: delineate causality between AEWs and TCs

# Seasonal basin-wide Atlantic TC numbers can be maintained in the absence of AEWs (for an active hurricane season), suggesting:

- ✤ Atlantic TCs are not limited by AEWs on season climate timescales.
  - ✤ AEWs are important for TC genesis on synoptic scale.
  - Spatial pattern of TC tracks?
- Covariability in AEWs and TCs may be driven by ocean variability.
- Although TCs readily generate from AEWs, in the absence of AEWs TCs will generate by other mechanisms.
  - The specific type of disturbance is unimportant for determining basin-wide seasonal Atlantic TC number.
- AEWs are not a reliable predictor of variability and change in *basin-wide* Atlantic TC frequency
  - ✤ … So what are predictability sources?





- Atlantic and Pacific SST variability jointly
  - Compensating and constructive influences on TCs
  - Extremely active/inactive hurricane seasons driven by ocean variability, not internal atmospheric variability.
- El Niño's spatial patterns
  - SST warming near warm pool is more effective than warming of cold tongue at suppressing Atlantic TCs.
- Coupled model SST biases substantially impact simulated TC activity → improvements needed for better prediction/projection. (Hsu et al. 2018)





# **ENSO Longitude Index (ELI)**



- Accounts for the nonlinear relationship between SST and deep convection.
- Calculation requires only monthly SST.
- Tracks the average longitude of tropical Pacific deep convection.

10-1

elative probability

10-3

-3 -2 -1 0

Nino3.4

 Characterizes the diversity and extremes of ENSO in a single index.

10

e probability

10-3

140 150 160

180 190 200 210 220

FI I



 ELI reveals increases in extreme El Niño, La Niña, and Modoki events at the expense of neutral ENSO in 21<sup>st</sup> century climate projections from CESM-LENS.

Williams and Patricola (2018) Diversity of ENSO Events Unified by Convective Threshold Sea Surface Temperature: A Nonlinear ENSO Index. GRL



Special thank you to:

- PIRATA organizing committee and participants
- sponsors IRD and Météo-France
- CLIVAR, NOAA, INPE

#### **U.S. Department of Energy**

Office of Science, Office of Biological and Environmental Research, Climate and Environmental Sciences Division, Regional & Global Climate Modeling Program (Award Number DE-AC02-05CH11231)

#### **National Science Foundation**

AGS-1347808 and OCE-1334707

#### **Texas Advanced Computing Center (TACC)**

at the University of Texas at Austin, in conjunction with the Extreme Science and Engineering Discovery Environment

### **Texas A&M Supercomputing Center**







