

# Vertical turbulent cooling of the mixed layer in the tropical Atlantic ITCZ and trade wind regions

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# Questions

Is there significant turbulent cooling at off-equatorial locations?  
Does it vary seasonally?

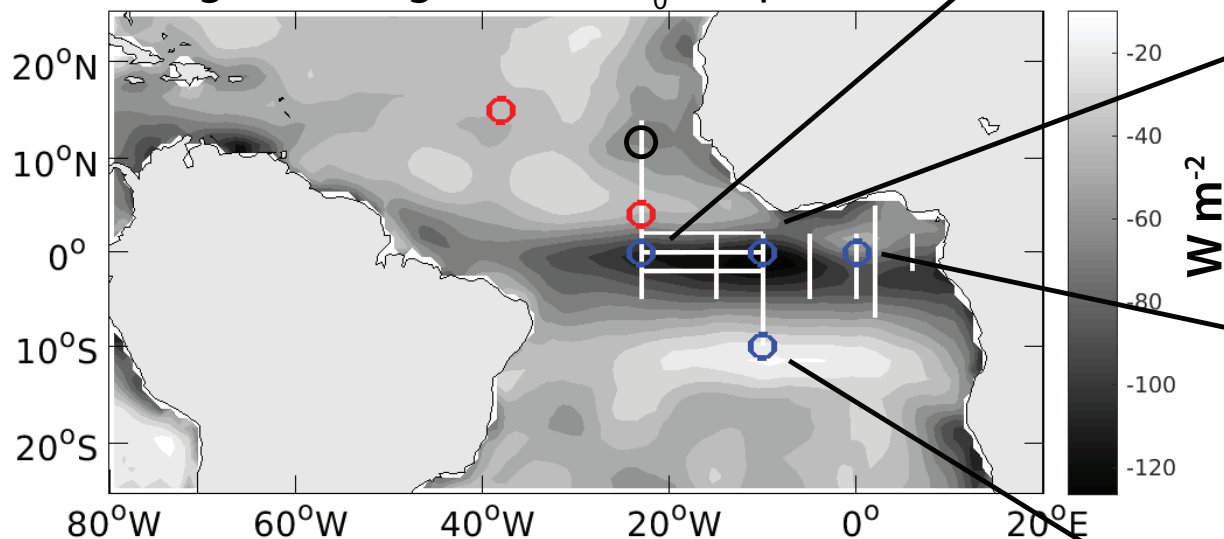
In the absence of strong mean shear, what drives mixing and cooling off the equator?

# Approach

Hourly measurements from PIRATA moorings and one-dimensional models (KPP, PWP).

# Previous measurements and estimates of turbulent cooling

**Shading:** annual mean heat budget residual:  
Storage rate (Argo) minus  $Q_0$  (TropFlux)

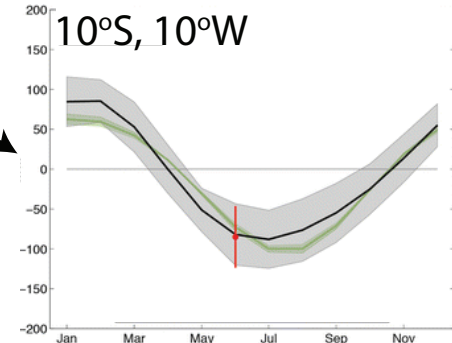
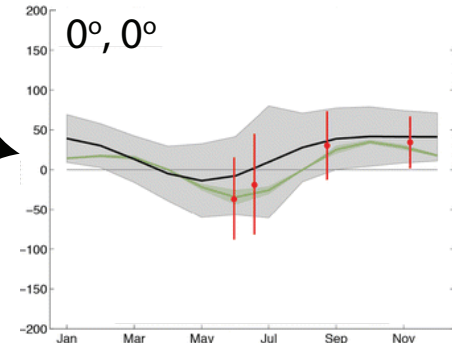
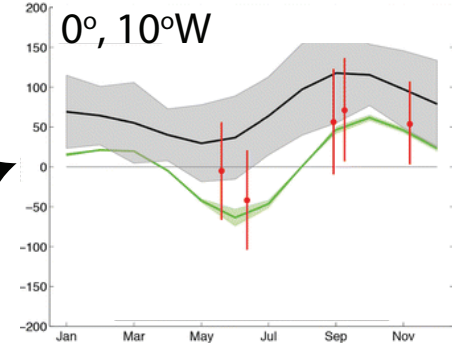
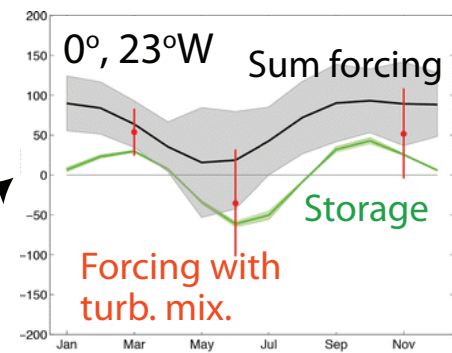


White lines: microstructure measurements

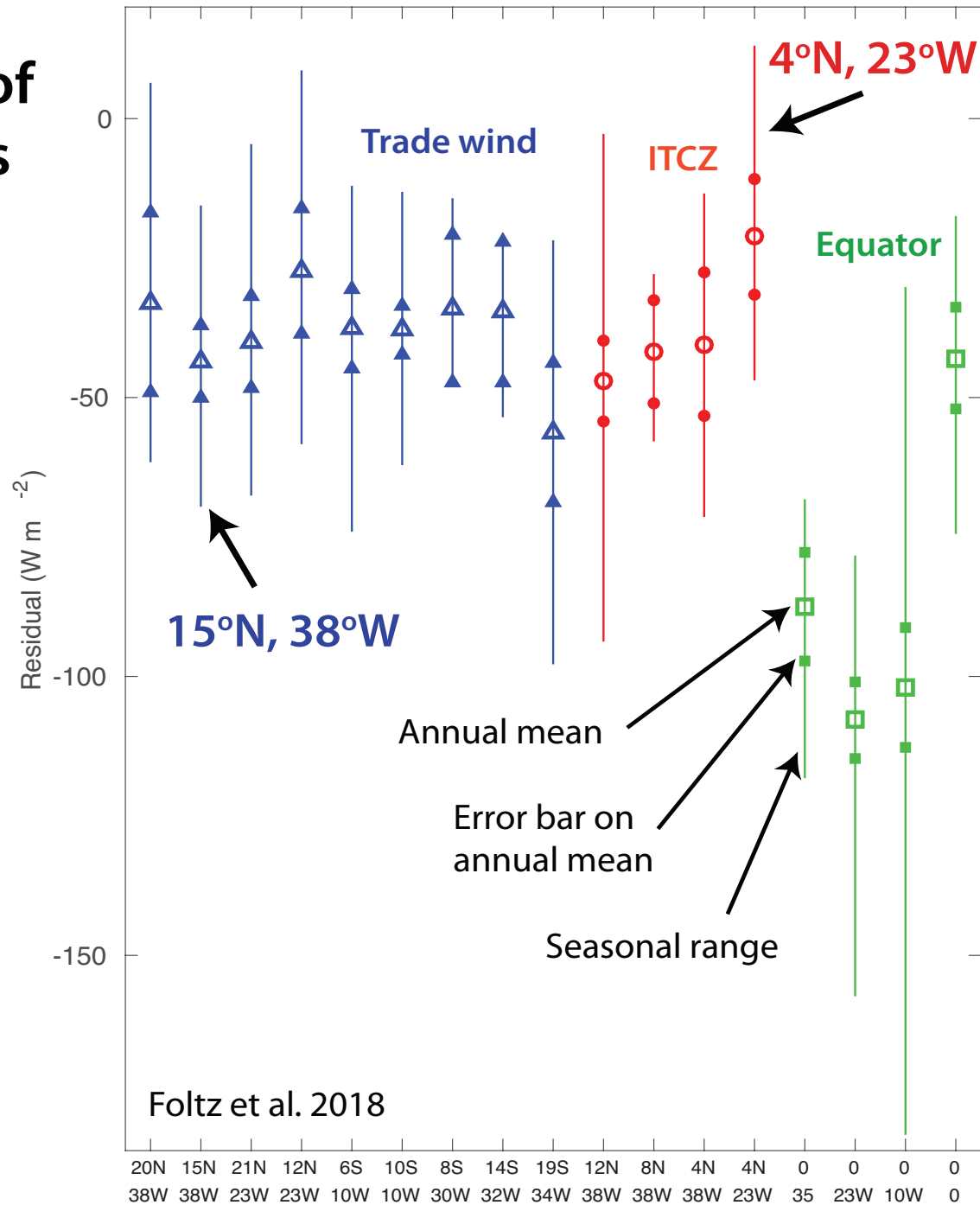
Blue circles: Hummels et al. (2014)

Black circle: Hummels et al. (2018)

Red circles: this study



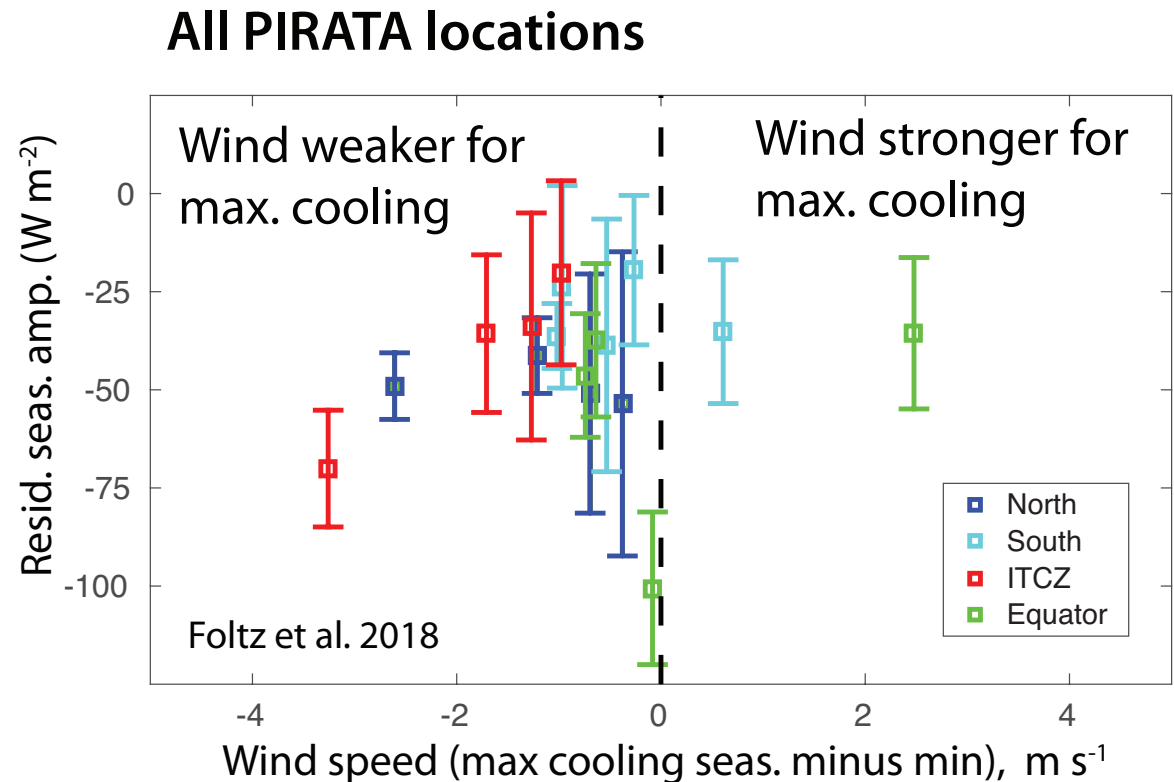
# Strong seasonalities of heat budget residuals





# Relationship between residual and wind speed

- Seasonally, more cooling occurs when wind is weak



# Data and methods

## 4°N, 23°W

- Hourly PIRATA **temp.** (1, 10, 20, 40, 60, 80, 100, 120, 140, 180 m)  
**salin.** (1, 10, 20, 40, 60, 120 m)  
**vel.** (7, 12, 17, 22, 27, 32, 37, 47, 57, 67, 87 m)  
**air temp., rel. humidity, winds, shortwave, rain**
- March 2017 - March 2018
- Calculate vertical diffusivity ( $K_v$ ) using KPP model (Large et al. 1994).

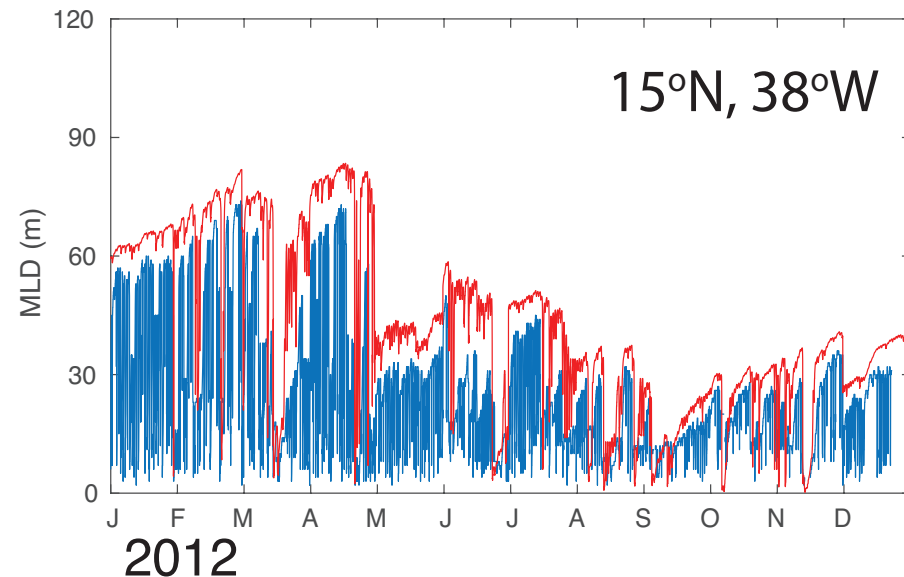
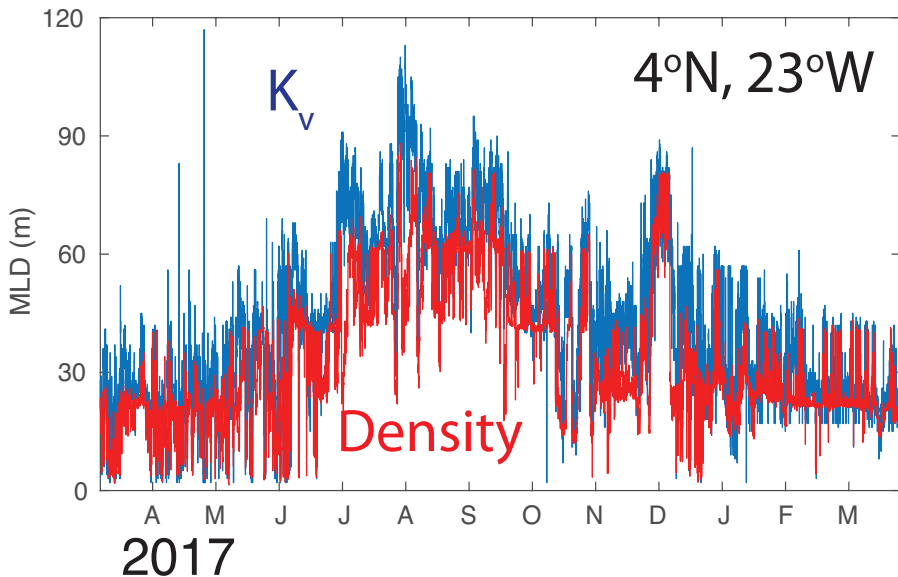
## 15°N, 38°W

- Daily ePIRATA **temp., salin.** (Foltz et al. 2018)  
Hourly PIRATA **air temp., rel. humidity, winds, shortwave, rain**
- Initialize **PWP model** (Price et al. 1986) at beginning of each month with ePIRATA  $T(z)$ ,  $S(z)$  then force with hourly winds, fluxes (2001, 2002, 2003, 2004, 2006, 2007, 2012). **84 monthly model runs.**
- Calculate  $K_v$  using KPP model: mooring sfc. forcing, PWP  $T(z)$ ,  $S(z)$ ,  $v(z)$ .

# Methods

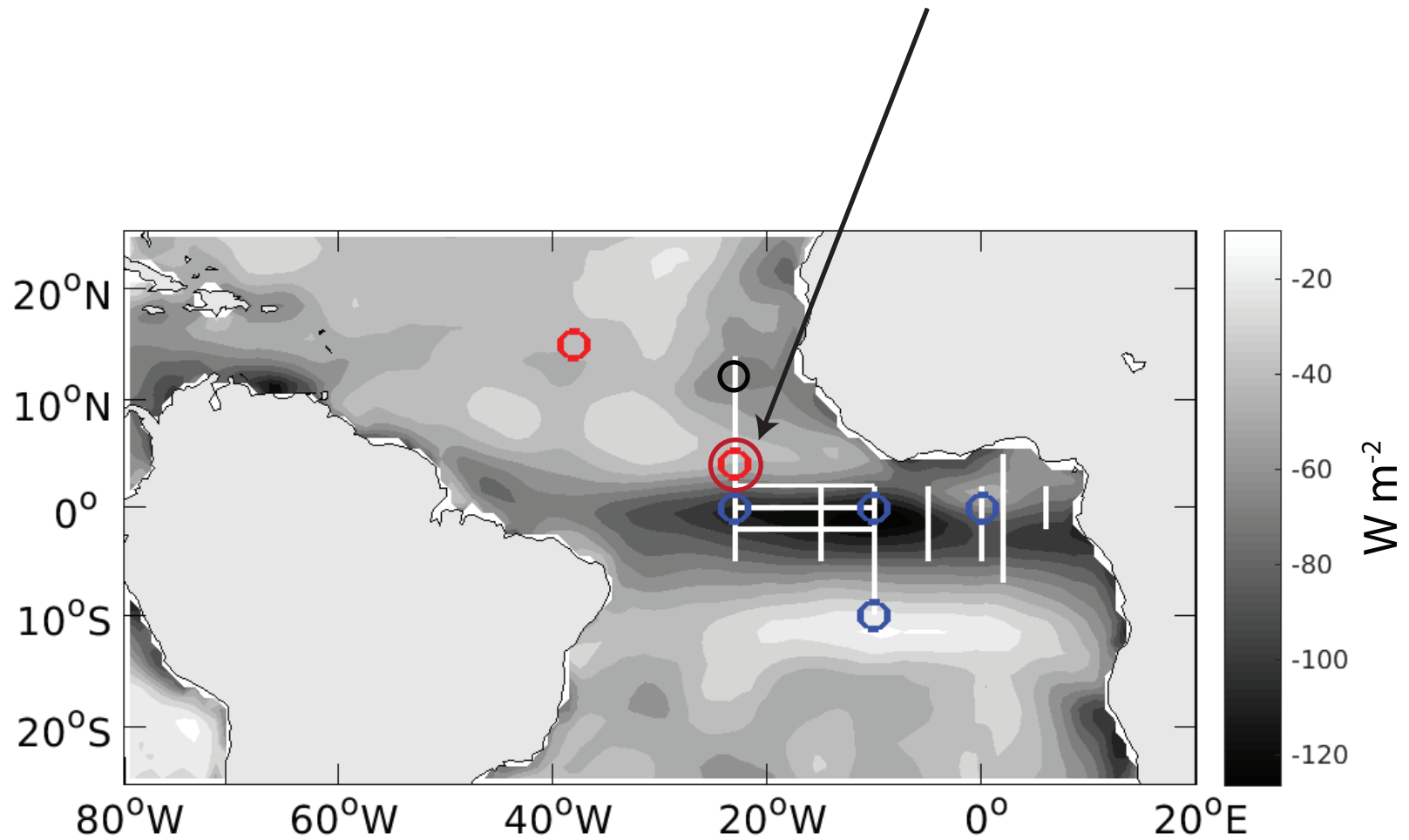
**MLD:** shallowest depth where  $K_v$  is less than  $0.001 \text{ m}^2 \text{ s}^{-1}$

## Comparison of $K_v$ -based and density-based MLD



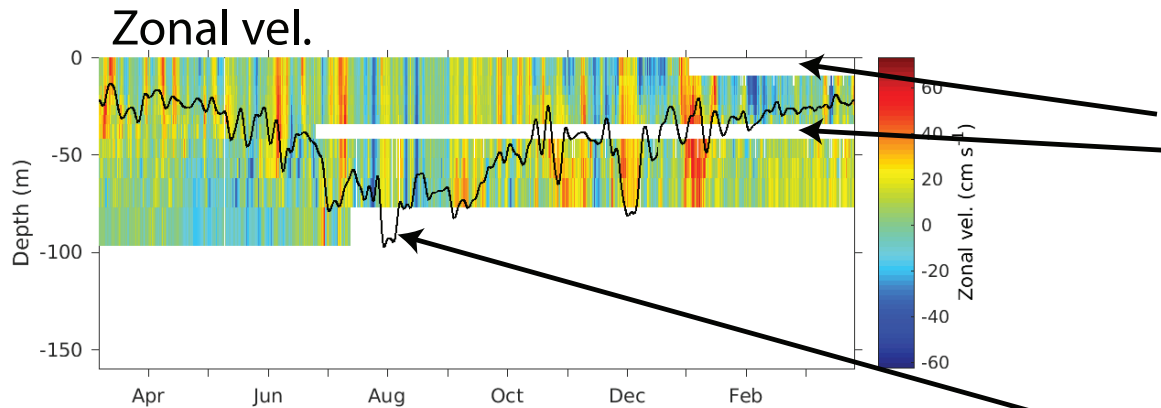
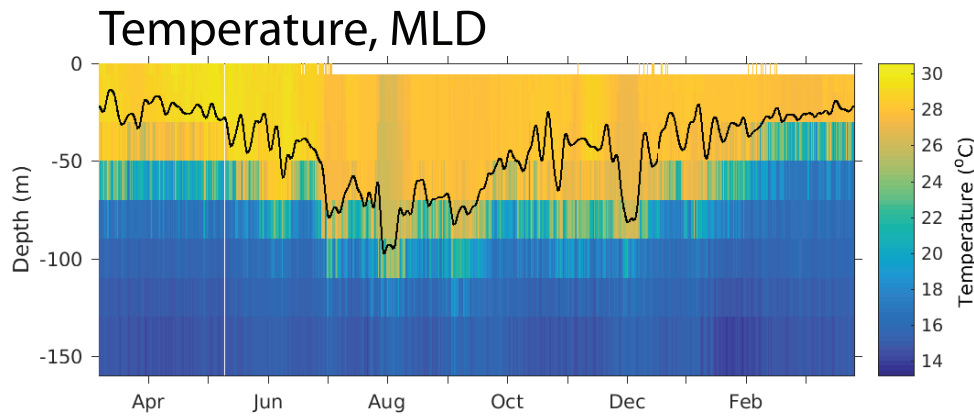
**Vert. turb. cooling of ML:**  $K_v$  at  $\text{MLD}+10 \text{ m}$  and  $dT/dz$  calculated between  $\text{MLD}$  and  $\text{MLD}+10 \text{ m}$ :  $\text{dens} * c_p * K_v * dT/dz$

# Results at 4°N, 23°W

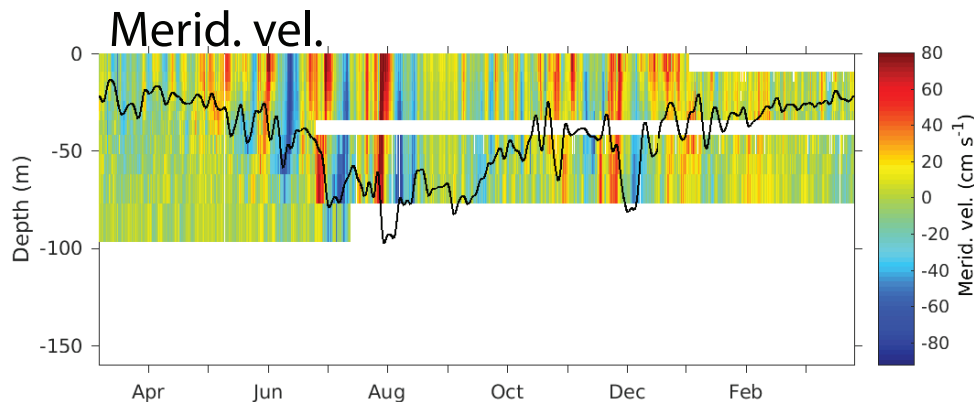




# Data from the 4°N, 23°W mooring



Fill with linear interpolation

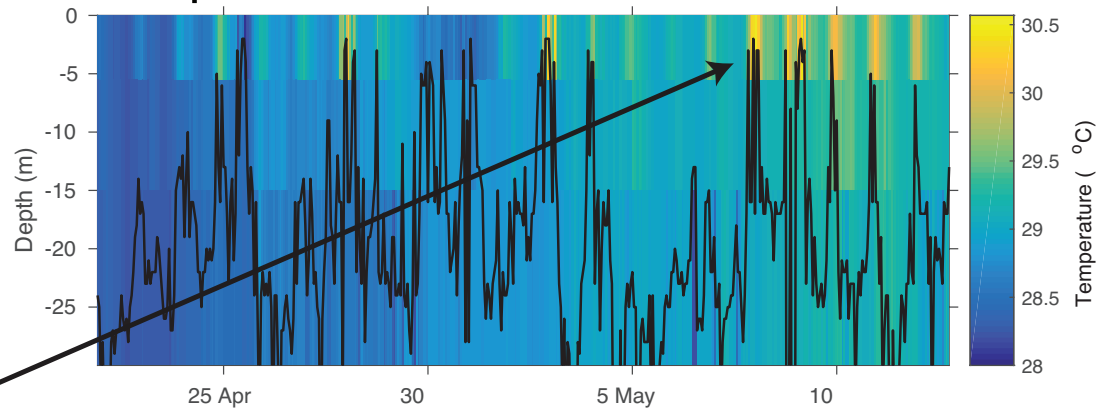


Leave gaps when MLD is within 10 m of deepest vel. measurement

2017

# Temp., vel. during April-May

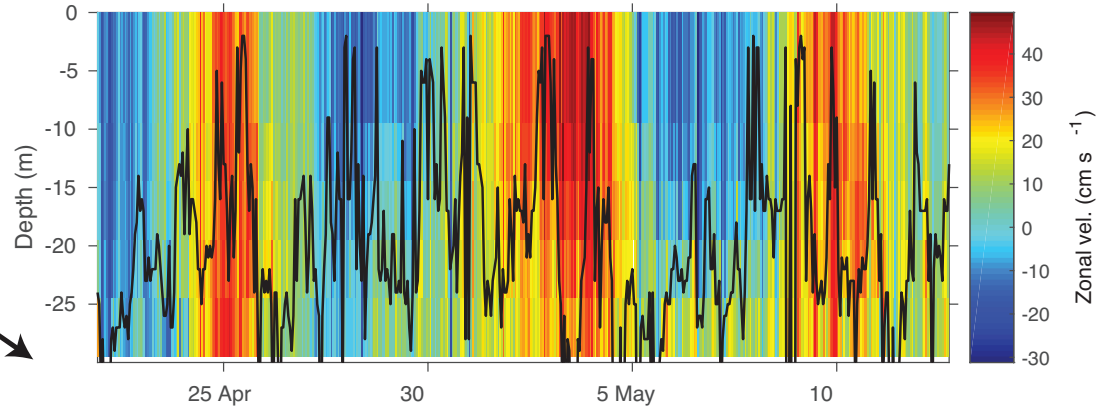
## Temperature, MLD



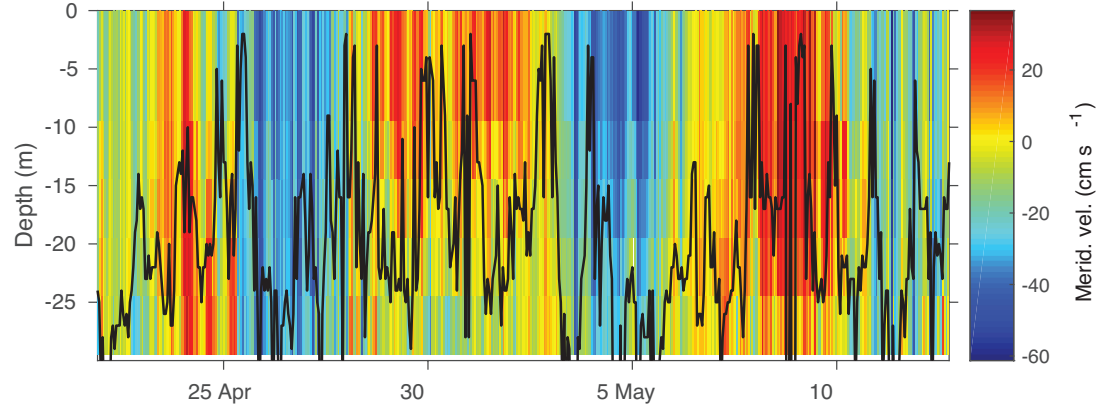
Diurnal temp.

Near-inertial waves,  
semidiurnal velocity

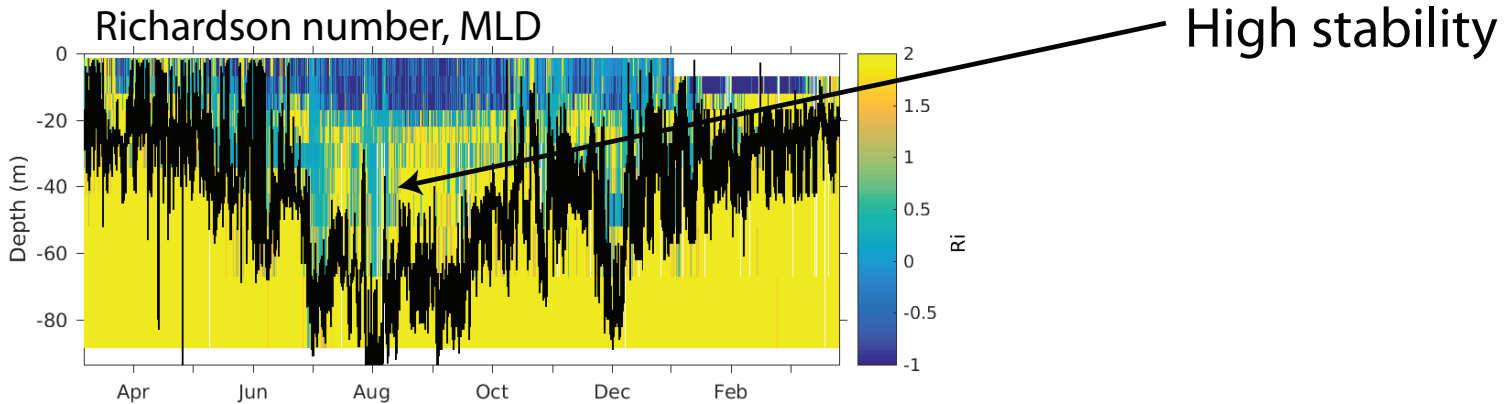
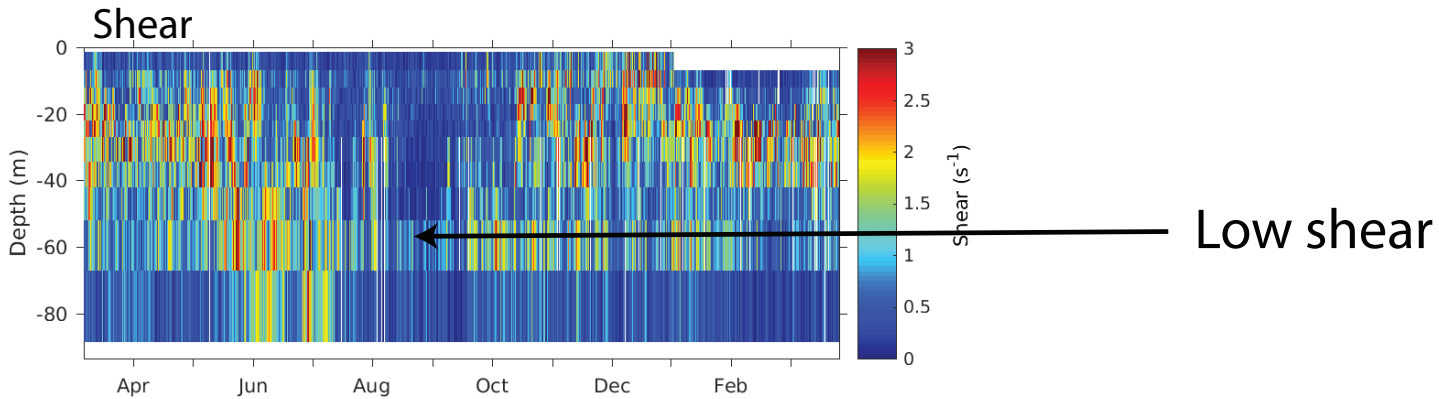
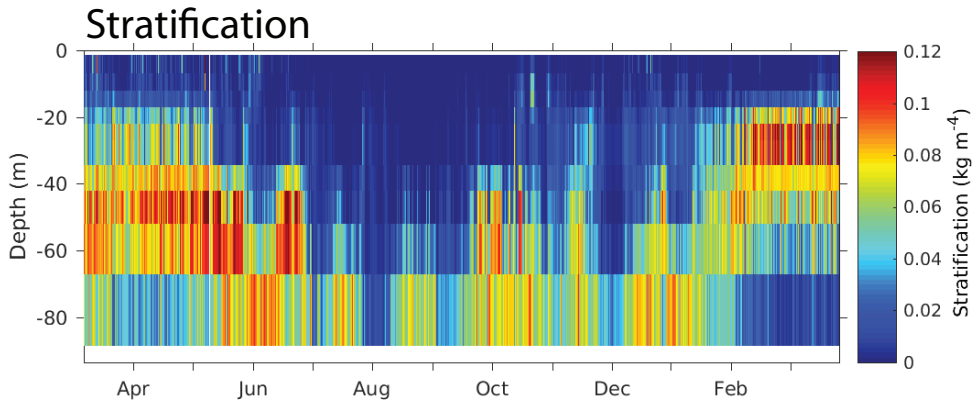
## Zonal vel.



## Merid. vel.

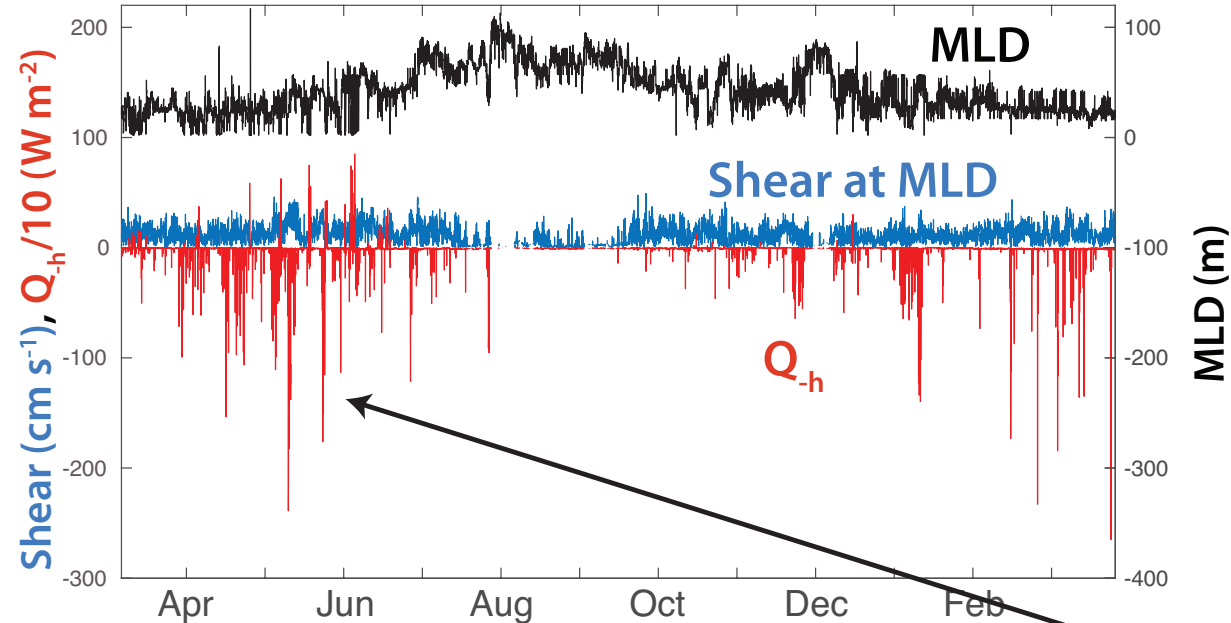


# Strong seas. cycles of stratification, shear

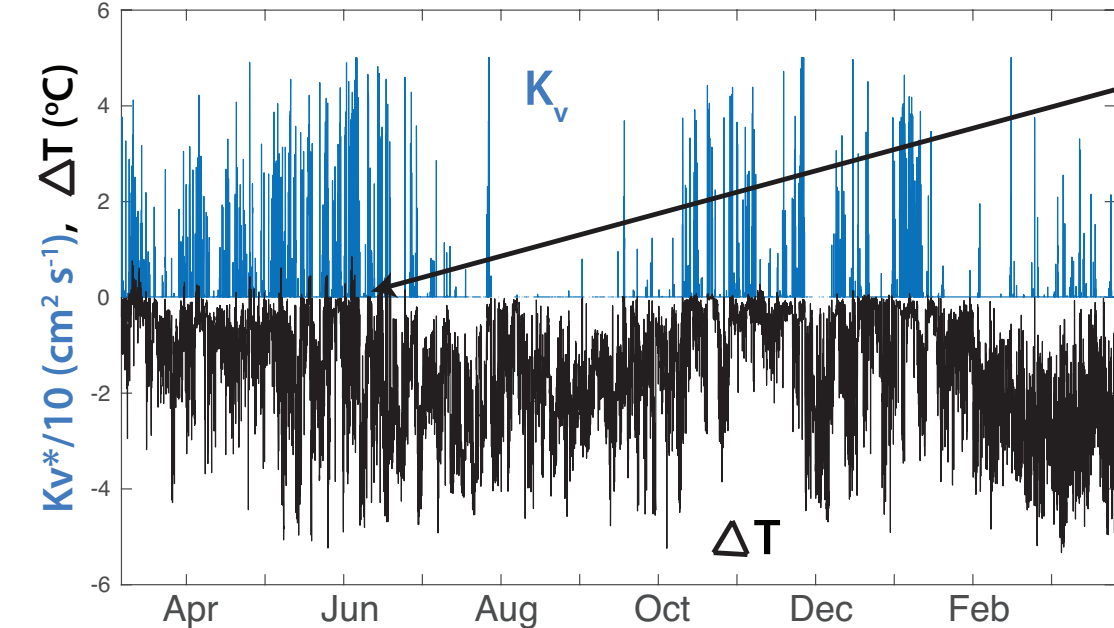


2017

# Vertical mixing and cooling



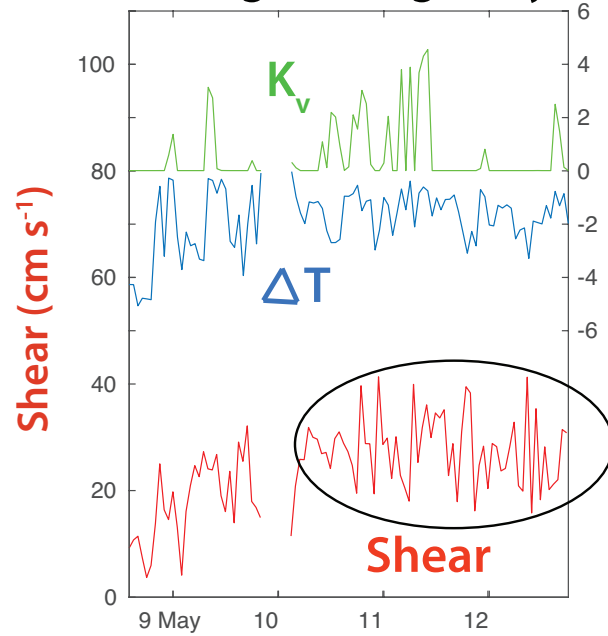
Largest  $K_v$ , strongest cooling when shear is strong, stratification ( $\Delta T$ ) is weak



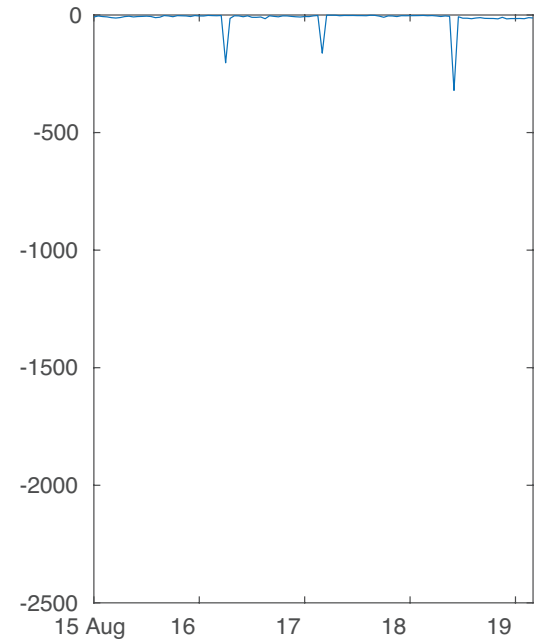
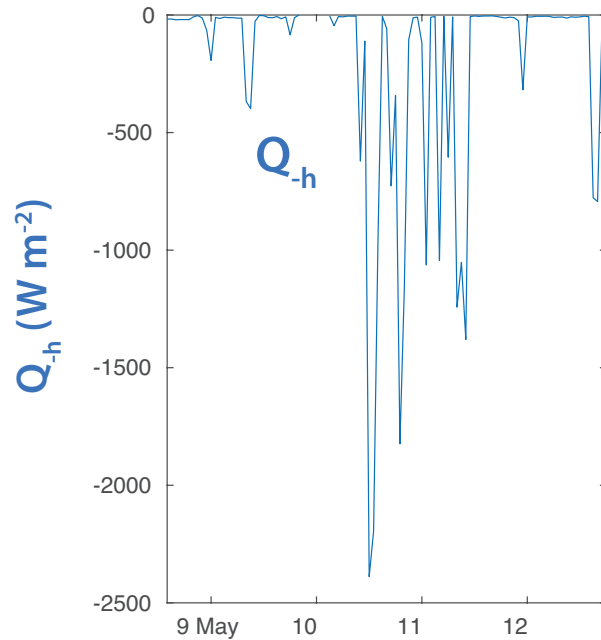
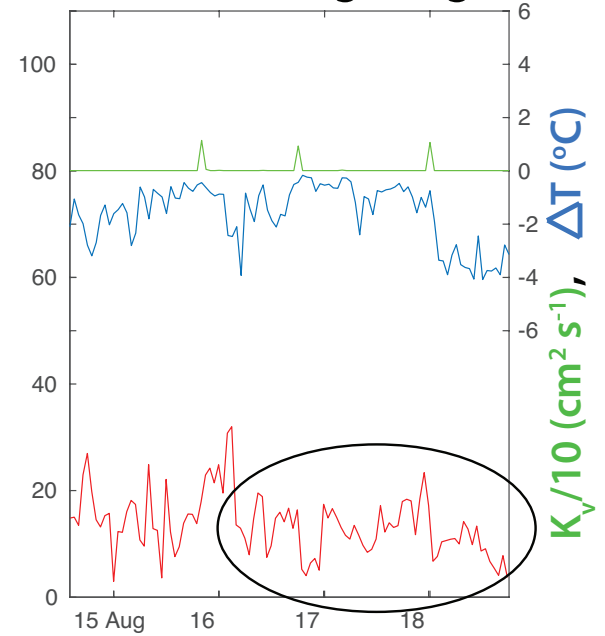


# Importance of shear and stratification

### Strong cooling (May)

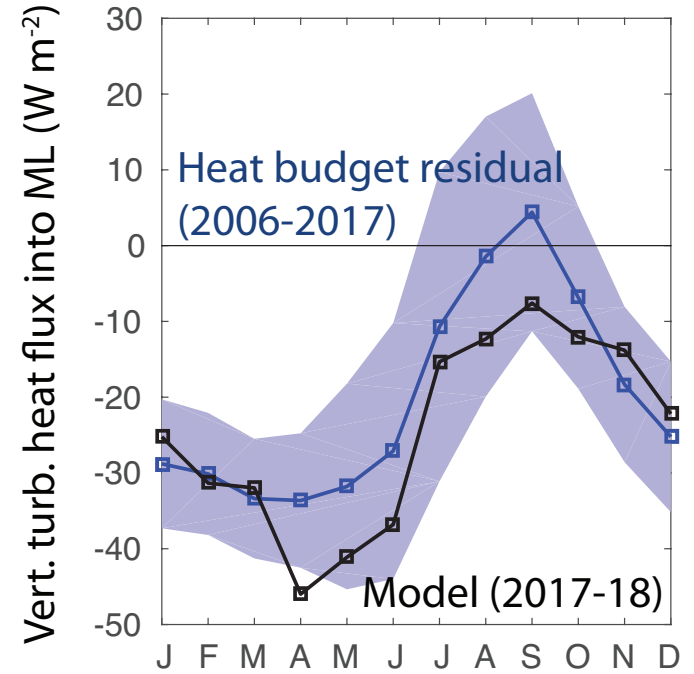


### Weak cooling (Aug)

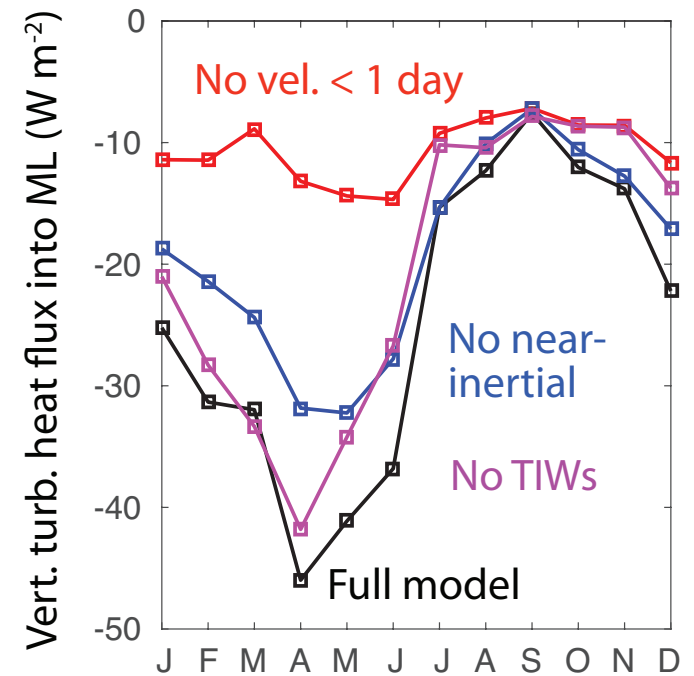


# Validation and diagnosis

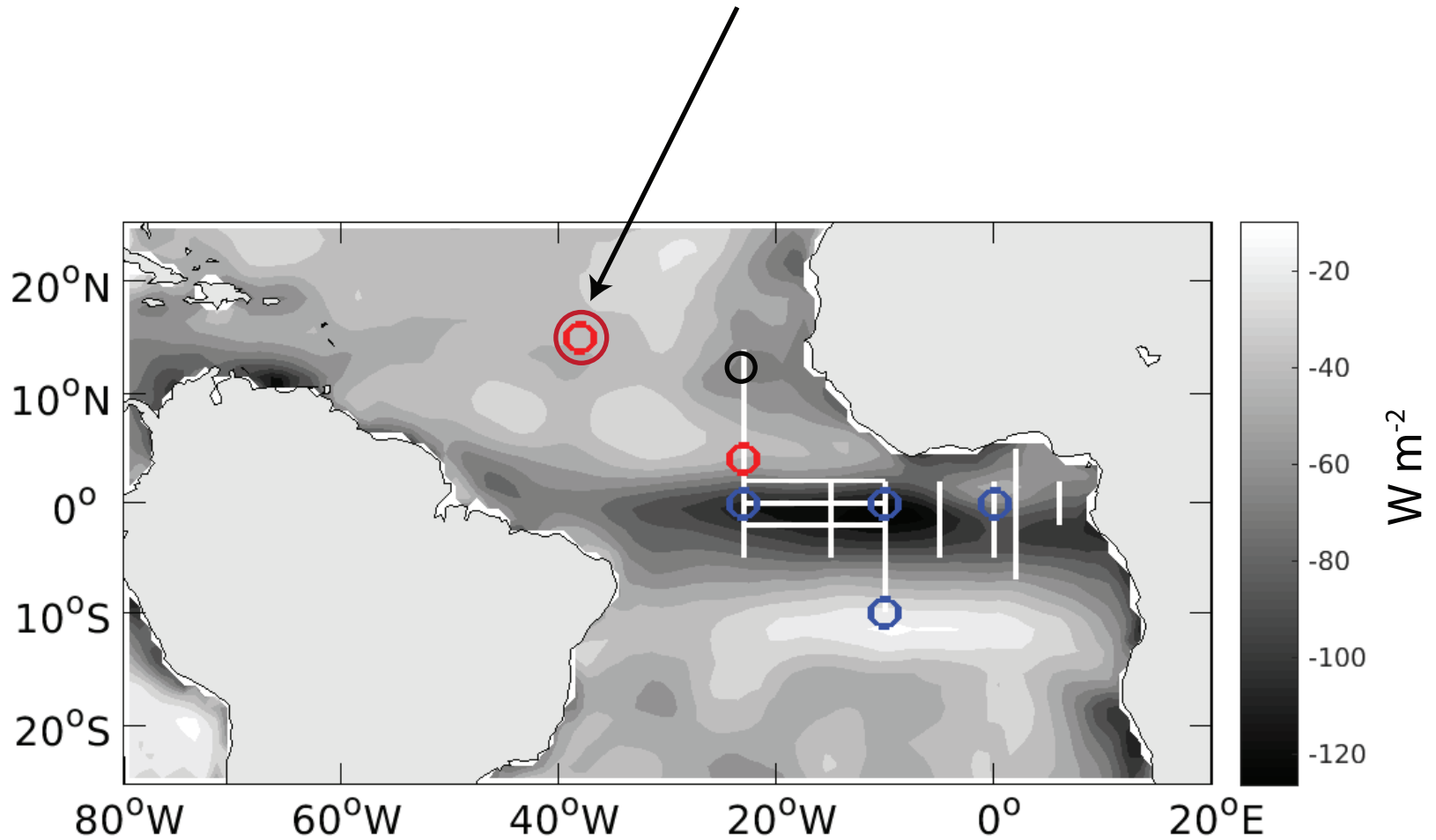
Seasonal cycle from model agrees well with that of heat budget resid.



Short timescale variability of shear (periods  $< 1$  day) is most important, likely due to tides and remotely-forced internal waves. **Near-inertial waves enhance cooling** during winter-spring.



# Results at 15°N, 38°W

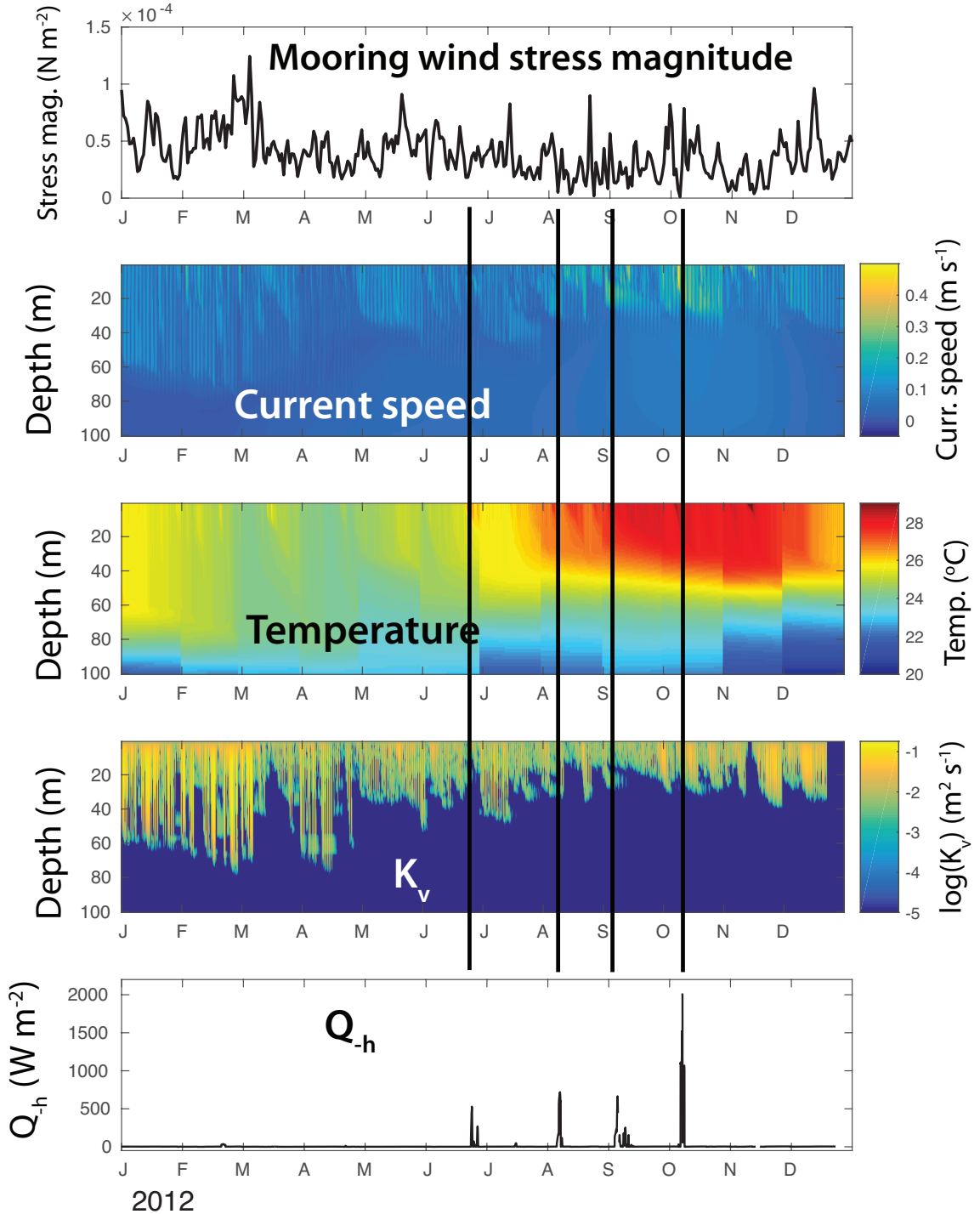


# Results from 2012 (PWP, KPP)

Strongest cooling occurs  
during summer-fall:

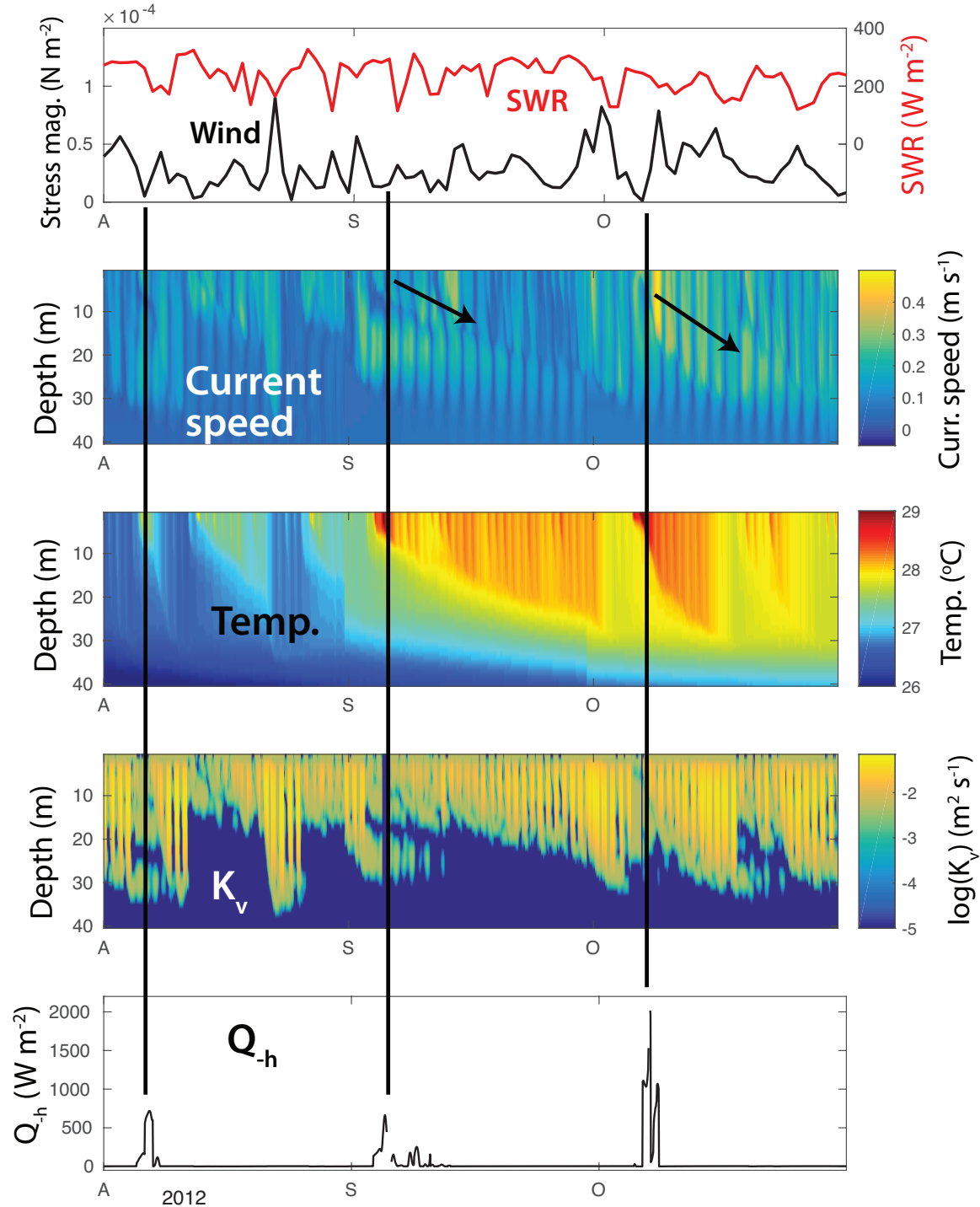
First: weak wind,  
surface warming,

Then: stronger wind,  
temp. and currents  
mixed downward,  
episodic ML cooling

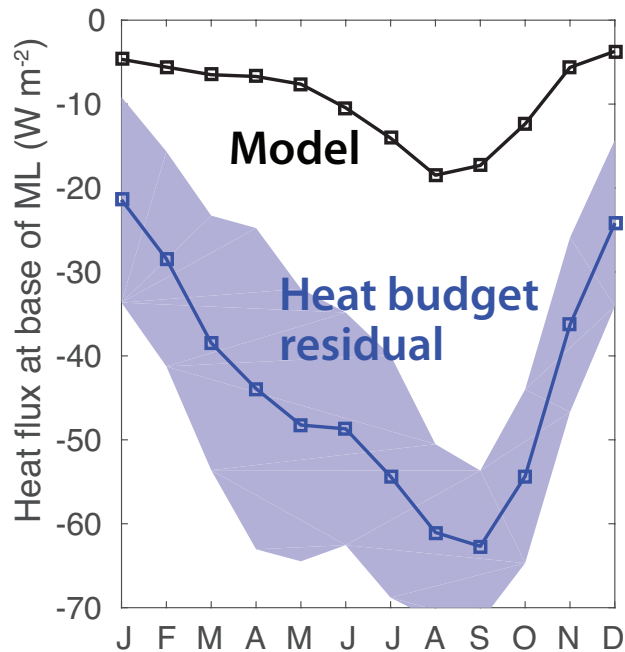




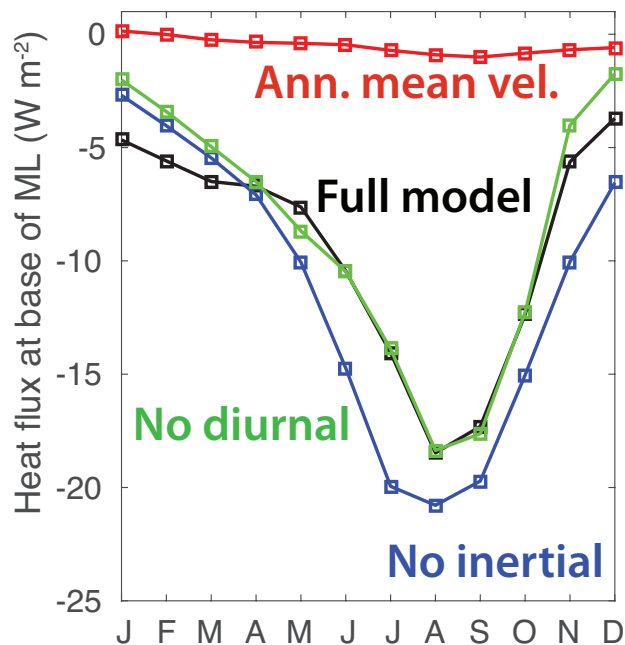
# Aug-Oct 2012



# Means seas. cycle (7 years, 2001-2012)



Phases of seas. cycles are similar, but model underestimates cooling, likely because tides and internal waves are missing.



Turbulent mixing in model is driven mainly by episodic shear. Weak influence from diurnal cycle and near-inertial waves.

# Summary and conclusions

- There are pronounced seasonal cycles of turbulent cooling at off-equatorial locations.
- Cooling tends to be strongest when winds are weakest and the mixed layer is thinnest. These conditions lead to enhanced shear at the base of the ML, which appears to originate mainly from remotely-forced internal waves with periods  $< 1$  day.
- Local wind- and buoyancy-forced mixing accounts for at most  $\sim 25\%$  of the seasonal cycle of cooling.
- Many unanswered questions remain, including the sources of remotely-forced shear and turbulence.
- These results need verification from direct measurements of turbulence.

# Validation of PWP model at 15°N, 38°W

